

Mon Oct 20

## Announcements:

- **bring calculator** to class from now on
  - > in-class activities
  - > midterm and final
- 

## Today:

- radiation and temperature (cont)
- sun-earth geometry
- energy balance >> conceptual model of climate change

Tues:  $T_e$  (as opposed to  $T_s$ ) and the greenhouse effect  
energy budget, balance at different levels

Wed: atmospheric composition and structure  
greenhouse gas chemistry (why do these molecules absorb IR?)

Thurs: multiple roles of clouds in energy balance  
local energy balance

## upcoming talks

Monday 20 October:

**3:30 310 ATG**

Prof Jim Tillman, UW

Climate Mars: Viking Lander and next generation

TUESDAY 21 October

**12:30 ATG 310c**, Weather discussion, Rick Steed

FRIDAY 24 October:

**3:30 14 OTB (Oceanography Teaching Bldg)**

Dr. Ralph Keeling

Scripps Institute of Oceanography

"Fate of anthropogenic CO<sub>2</sub> and changing biogeochemistry  
of the oceans"

## Radiation/Wavelength note

### 1. Electromagnetic radiation:

- propagates through space
- has a characteristic **wavelength**,

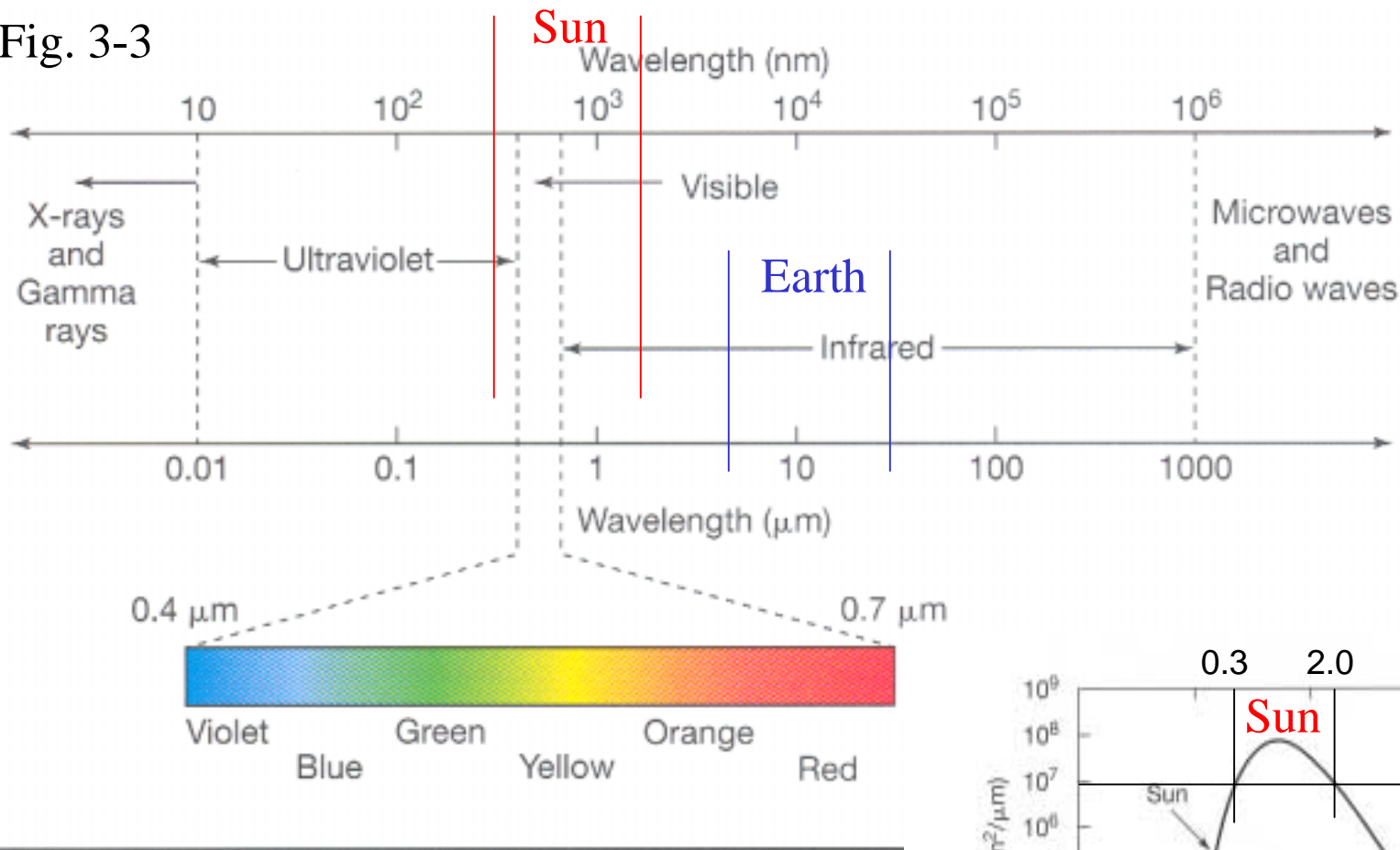
for now just two types:

**SW: shortwave** (mostly visible)

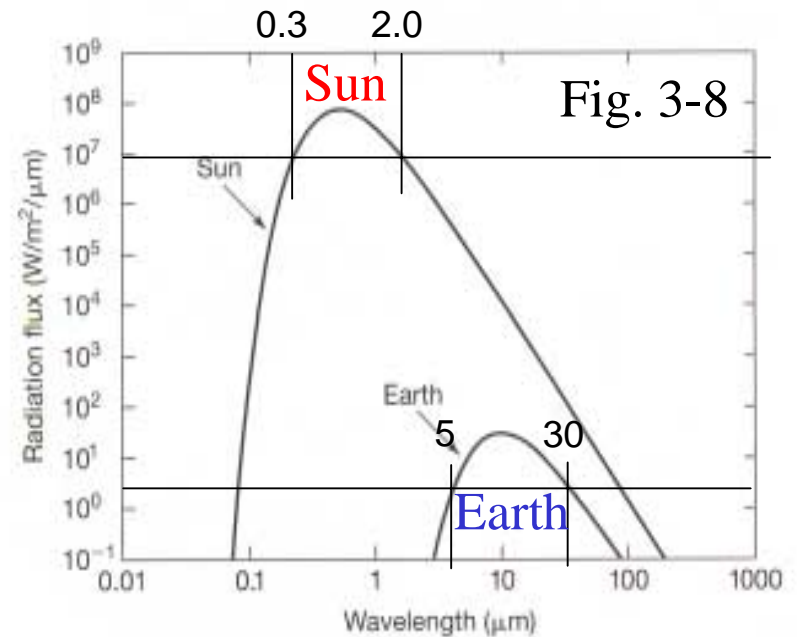
**LW: longwave**

SW and LW ... not same meaning as radio !

Fig. 3-3



Radiation vs  
Wavelength



# Stefan-Boltzmann Law

Hotter objects emit **more total energy**:

$$\text{Stephan-Boltzmann Law: } E = \sigma T^4$$

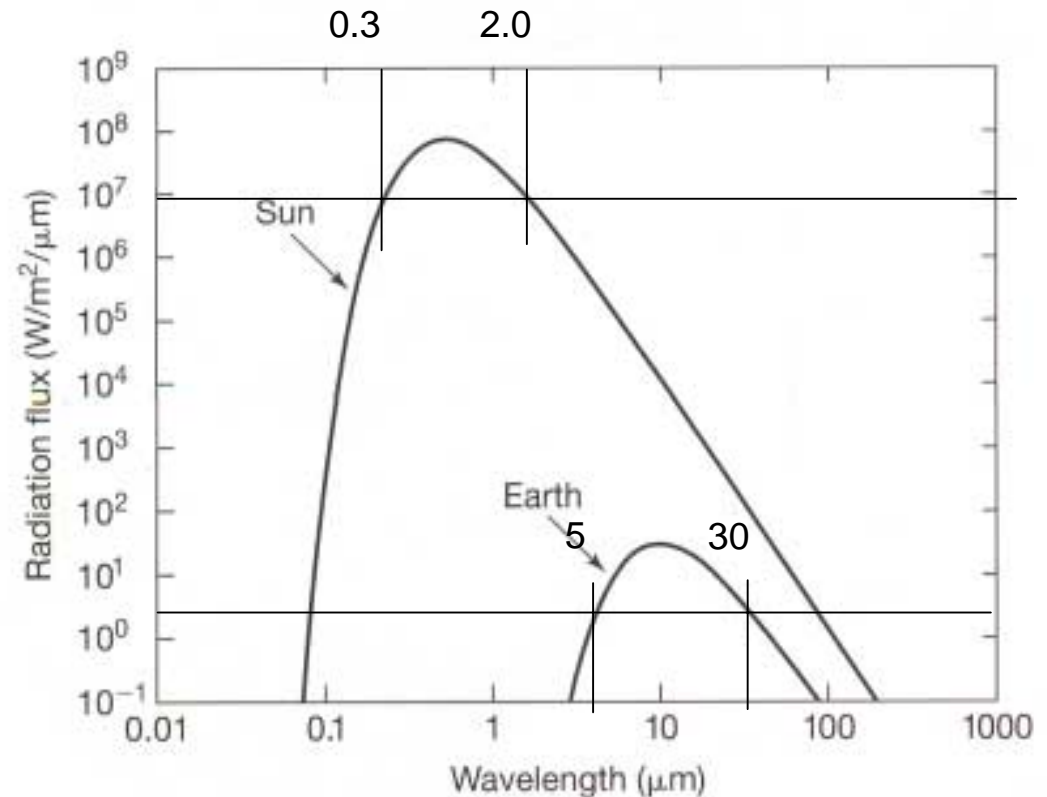
$$\sigma = 5.67e-8 \text{ [a constant]}$$

T is temp in Kelvin, the absolute temperature scale

$$T(K) = T(C) + 273$$

Applies to an ideal material known as a "blackbody".

(The emissions from most physical objects are close to this ideal.)



**FIGURE 3-8**

Blackbody emission curves for the Sun and Earth. The Sun emits more energy at all wavelengths.

## Kelvin (absolute) Temperature Scale

$$T(\text{K}) = T(\text{C}) + 273$$

Reference Point	Celcius Scale (C)	Kelvin Scale (K)
water boils	100	373
water freezes	0	273
????	-273	0

What is the meaning of "absolute zero" temperature?

Consider...

- Hotter objects emit more total energy
- Can there be a negative temperature on the Kelvin scale?
- What is the meaning of "temperature", anyway?

## Kelvin/Boltzmann applications

Stephan-Boltzmann Law:  $E = \sigma T^4$

where  $\sigma = 5.67e-8$ ,  $E$  is in  $W/m^2$ , and  $T$  is in Kelvin (K)

Kelvin vs Celcius scale:  $T(K) = T(C) + 273$

The average surface temperature of the Earth is 15C.

Express this in degrees Kelvin.

An object warms from 0C to 10C.

What is the fractional increase in temperature?

What is the fractional increase in emitted, radiative energy?

If the Sun were to warm up by 10%, by how much would the solar constant increase?

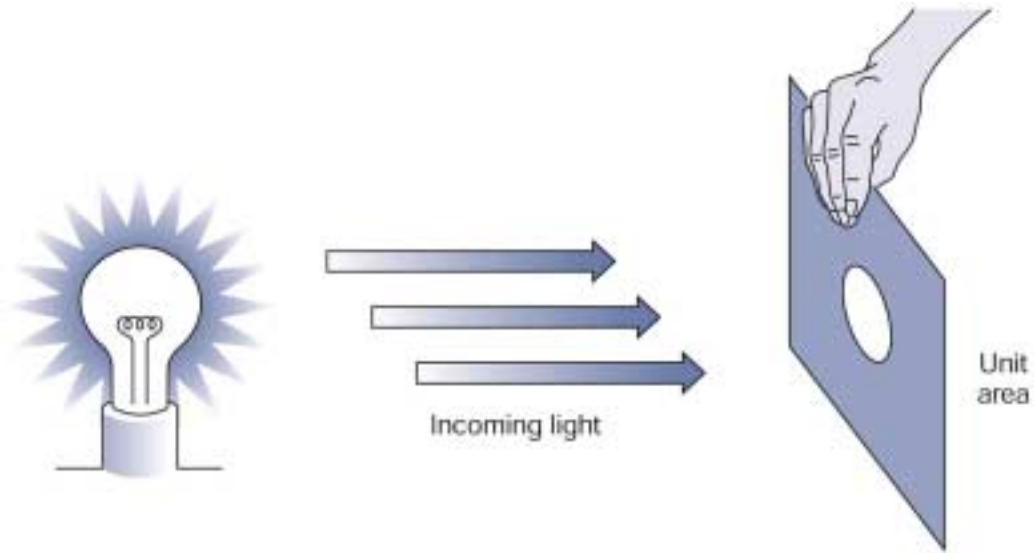
(i) fractional increase, (ii) absolute increase

$T_{SUN} \sim 6000$  K,  $T_{EARTH} \sim 255$  K

What is the ratio of energy flux,  $E$  ( $W/m^2$ ) from the Sun to that from the Earth?

# Sun-Earth Geometry: flux on an angled surface

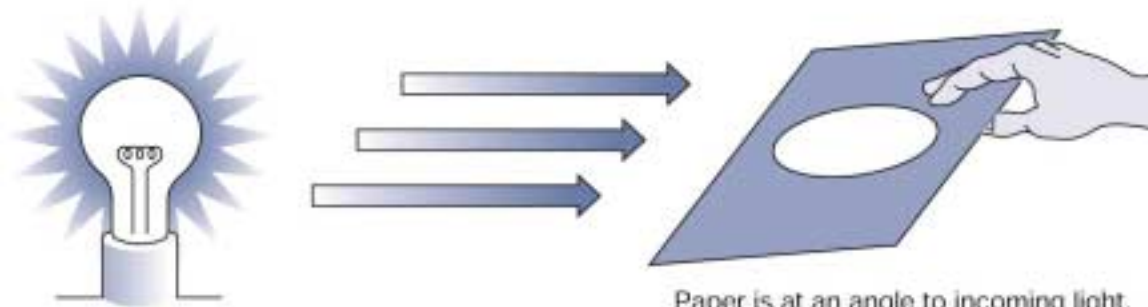
overhead sun,  
tropics



Paper is perpendicular to incoming light.

(a)

sun low in the sky,  
high latitudes



Paper is at an angle to incoming light.

(b)



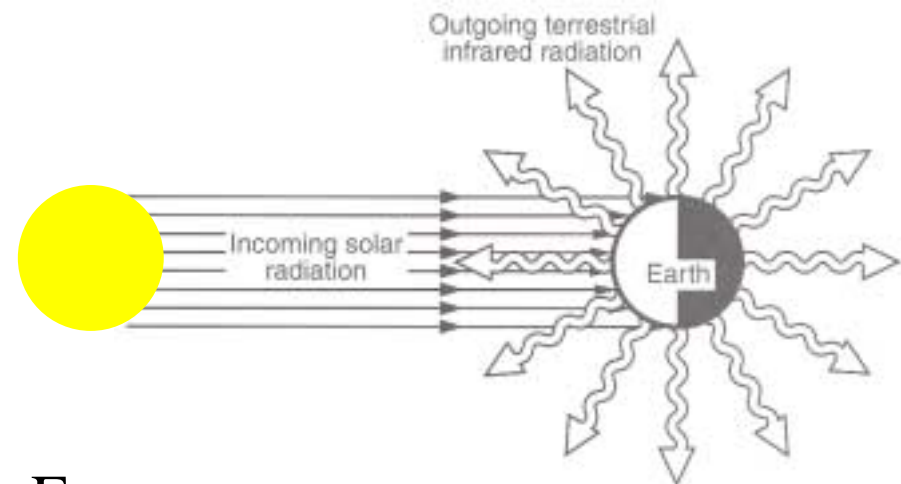
## Sun-Earth Geometry: factor of 4

Define the solar constant,  $S_0$ : the flux of solar energy passing through space at the Earth's orbital distance.

$$S_0 = 1370 \text{ W/m}^2$$

The earth intercepts the same amount of solar energy as flat disc of area  $\pi r^2$

The earth emits LW radiation over the surface area of a sphere,  $4\pi r^2$



$$E_{\text{IN}} = E_{\text{OUT}}$$

$$S_0(1-A) \pi r^2 = \sigma T_e^4 4\pi r^2$$

## Energy balance

$$E_{\text{IN}} = E_{\text{OUT}}$$

$$S_0(1-A) \pi r^2 = \sigma T_e^4 4\pi r^2$$

Now divide both sides by the surface area of the Earth,  $4\pi r^2$ .  
This will give the energy fluxes averaged over the Earth's surface.

$$\frac{S_0}{4} (1-A) = \sigma T_e^4$$

**In-class activity:**

$$S_0 = 1370 \text{ W/m}^2$$

the average albedo of the Earth is 0.30

>> calculate  $E_{\text{IN}}$  <<

$$E_{\text{IN}} = 240 \text{ W/m}^2$$

The average flux of solar energy absorbed by the Earth system.  
(averaged over entire surface, including day/night)

Tues Oct 21

## Announcements:

- **bring calculator** to class from now on
  - > in-class activities
  - > midterm and final
- 

## Today:

- conceptual model of climate change: forcing > feedback > response
- "effective radiating temperature",  $T_e$
- greenhouse effect (property of a planet's atmosphere)
- energy budget: needs to balance at every level
- [everything you need for HW#2 by end of today]

# Energy Balance and Climate Change

Energy balance equation (E is energy flux, W/m<sup>2</sup>)...

$$\begin{array}{ccc} E_{\text{IN}} & = & E_{\text{OUT}} \\ \text{energy absorbed} & & \text{energy emitted} \\ \text{by Earth} & & \text{by Earth} \end{array}$$

## “Energy Balance Theory of Climate Change”

if  $E_{\text{OUT}} < E_{\text{IN}}$  [imbalance]

then  $T_{\text{surface}}$  will go up until balance is restored

## Energy Balance Theory of Climate Change

$$\Delta T_s = \lambda \Delta F$$

$\Delta F$ : forcing, change in energy balance ( $\text{W}/\text{m}^2$ )

$\Delta T_s$ : response, change in surface temperature (K)

$\lambda$ : feedbacks, climate sensitivity  $\{\text{K}/(\text{W}/\text{m}^2)\}$

What are the characteristics of each of these terms?

$\Delta F$ : imposed changes on the climate system;  
increasing GHGs, changed surface albedo, increasing aerosols

$\Delta T_s$ : climate response in terms of GAAST;  
can be measured and predicted for the past (test of understanding)

$\lambda$ : involves full complexity of the climate system; requires  
climate models; involves feedbacks, mainly (i) water vapor,  
(ii) ice-albedo, (iii) clouds

# Energy Balance Theory of Climate Change

$$\Delta T_s = \lambda \Delta F$$

$\Delta F$ : forcing, change in energy balance ( $\text{W}/\text{m}^2$ )

$\Delta T_s$ : response, change in surface temperature (K)

$\lambda$ : feedbacks, climate sensitivity  $\{\text{K}/(\text{W}/\text{m}^2)\}$

There are two important terms left out of the above equation.

What are they?

Response time of the climate system

Natural variability of the climate system

(We will incorporate these later.)

## Effective Radiating Temperature -1

$$E_{\text{IN}} = E_{\text{OUT}}$$

$$\frac{S_0}{4} (1-A) = \sigma T_e^4$$

where  $S_0$  is the solar constant

$A$  is the planetary albedo

$\sigma$  is the Stephan-Boltzmann constant

and  $T_e$  is the "effective radiating temperature"

Given values for  $S_0$ ,  $A$ , and  $\sigma$ , we can solve for  $T_e$ :

$$T_e^4 = \frac{S_0}{4\sigma} (1-A)$$

$$T_e = \left[ \frac{S_0}{4\sigma} (1-A) \right]^{0.25}$$

## Effective Radiative Temperature -2

$$T_e = \left[ \frac{S_0}{4\sigma} (1-A) \right]^{0.25}$$

$\sigma = 5.67\text{e-}8$  [universal constant]

For Earth:  $S_0 = 1370 \text{ W/m}^2$

$A = 0.30$

Calculate  $T_e$  ???

$$T_e = 255 \text{ K} \quad (\text{or } -18 \text{ C})$$

"The Earth radiates as if it were a blackbody at 255 K."

but...

the Earth's surface temperature ( $T_s$ ) is 288 K (or +15 C).

???



## Greenhouse Effect

If the Earth had no atmosphere (and still had an albedo of 0.30), its surface temperature would be 255 K.

The atmosphere acts like a blanket, trapping heat near the surface and keeping it much warmer than it would otherwise be.

The magnitude of this "greenhouse effect" ( $\Delta T_g$ ) is:

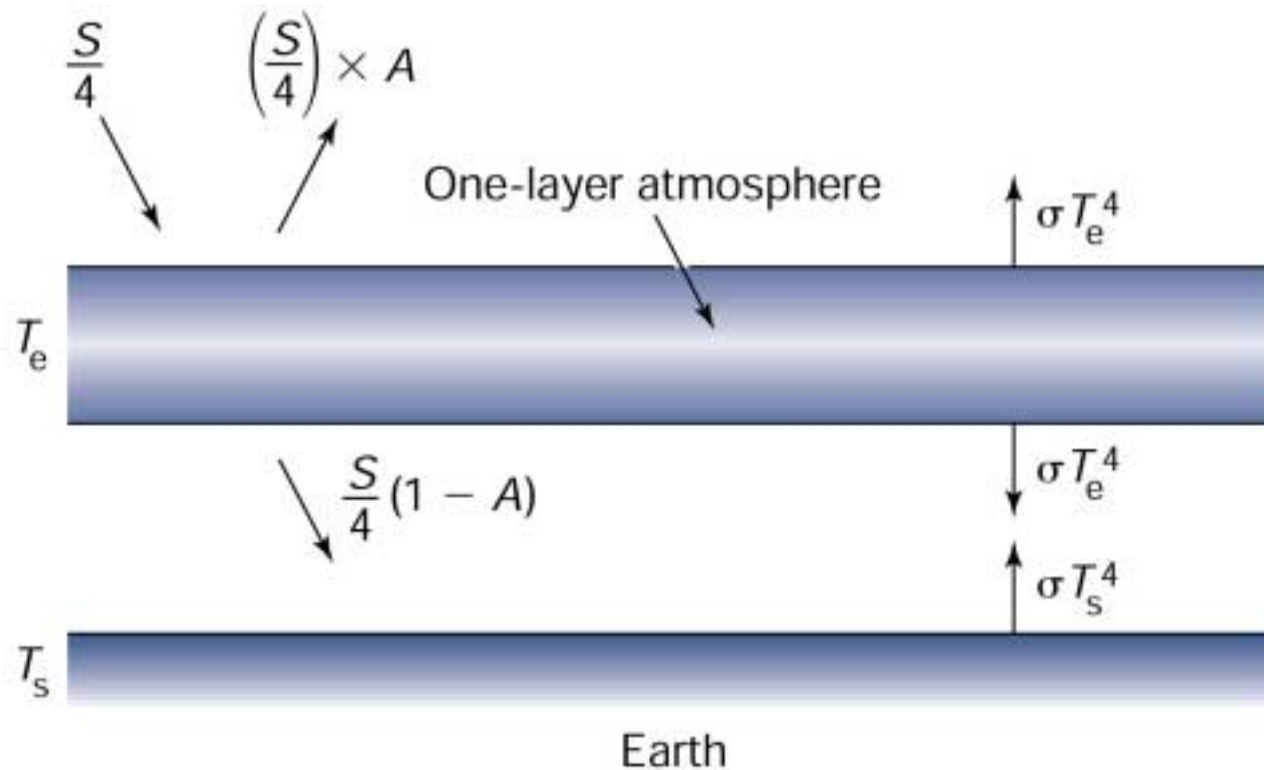
$$\Delta T_g = T_s - T_e = 288 \text{ K} - 255 \text{ K}$$

$$\Delta T_g = 33 \text{ K}$$

$\Delta T_g$  is a property of the atmosphere.

(Should really be called "the atmosphere effect".)

## Text Box on page 43



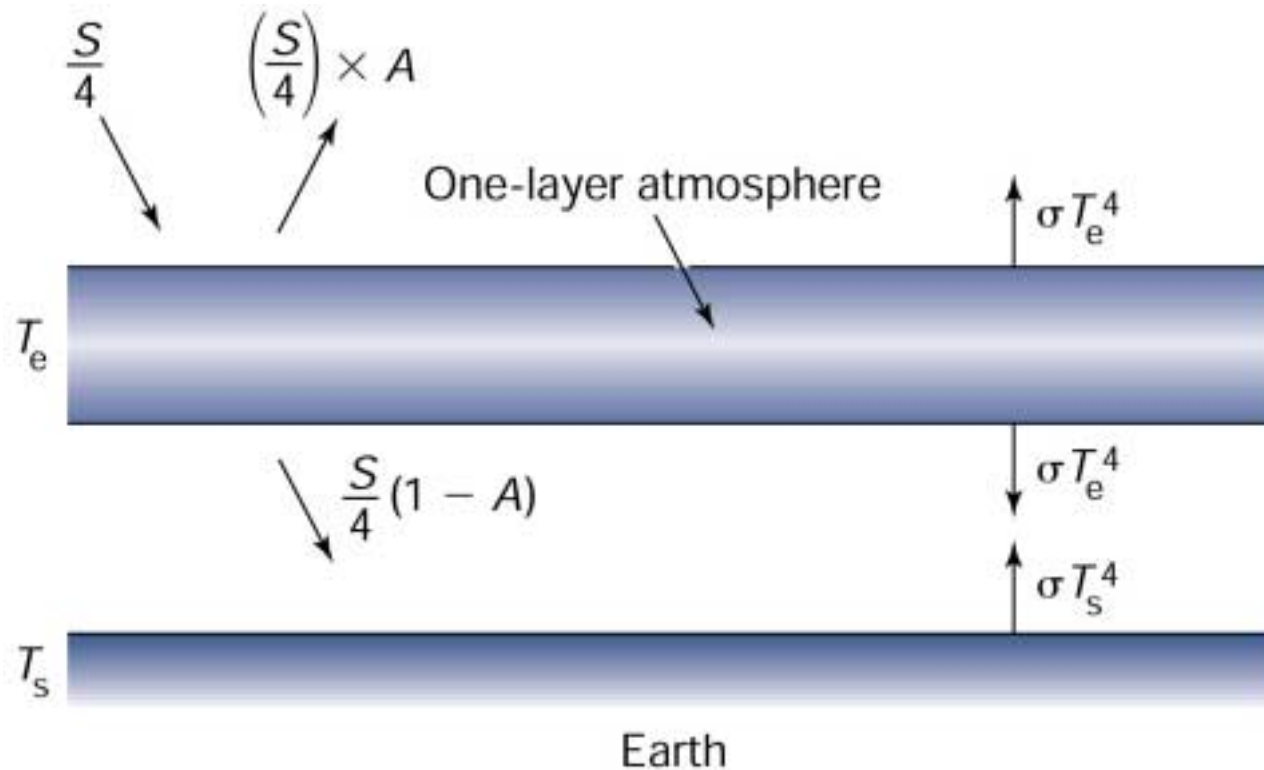
What are the assumptions of this model?

Atmosphere is a separate "layer" with temperature  $T_e$ .

Atmosphere does not absorb any SW radiation.

Atmosphere completely absorbs the LW radiation emitted by the surface.

## Text Box on page 43



**Good:** Atmosphere acts like a blanket.

Simple numerical model of the greenhouse effect.

**BAD:** Calculation is not very accurate ( $\Delta T_g = 48$  K)

Includes a misleading, incorrect equation:

$$\cancel{T_s = 2^{1/4} T_e}$$

## Sister planets

In fact, you cannot "predict"  $T_s$  from knowledge of  $T_e$ .

$T_e$  is a function of  $S_0$  and  $A$

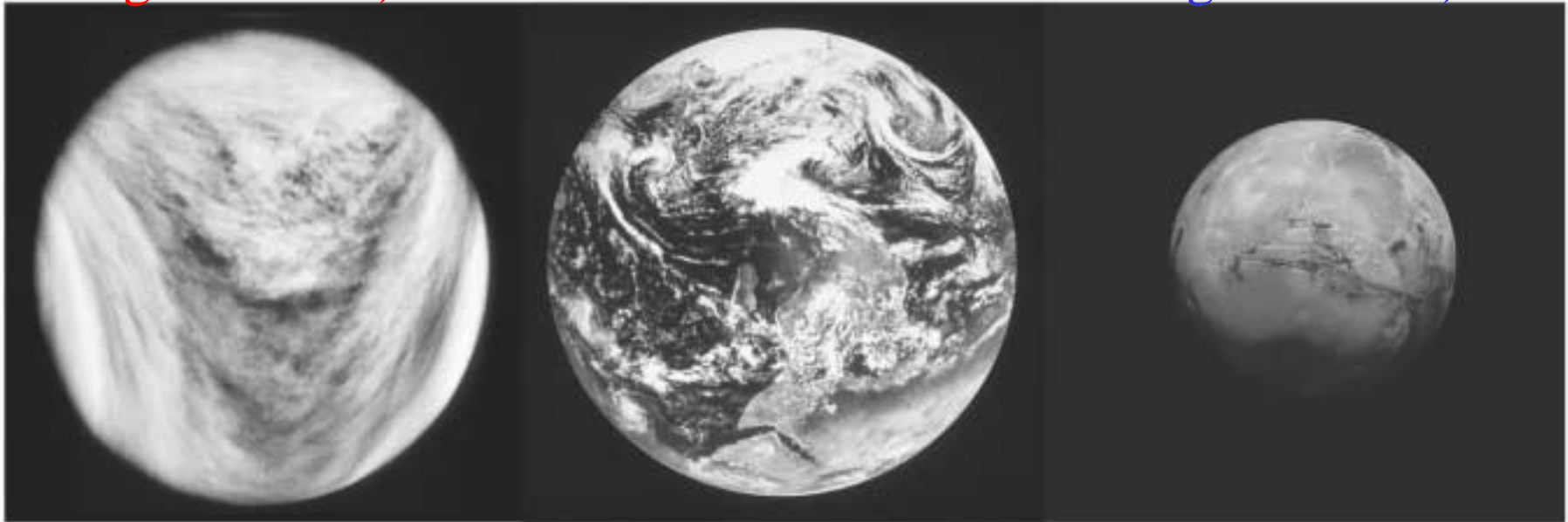
$$T_s = T_e + \Delta T_g$$

where,  $\Delta T_g$  is a property of the atmosphere of a planet

Venus  
(runaway  
greenhouse)

Earth  
("just right")

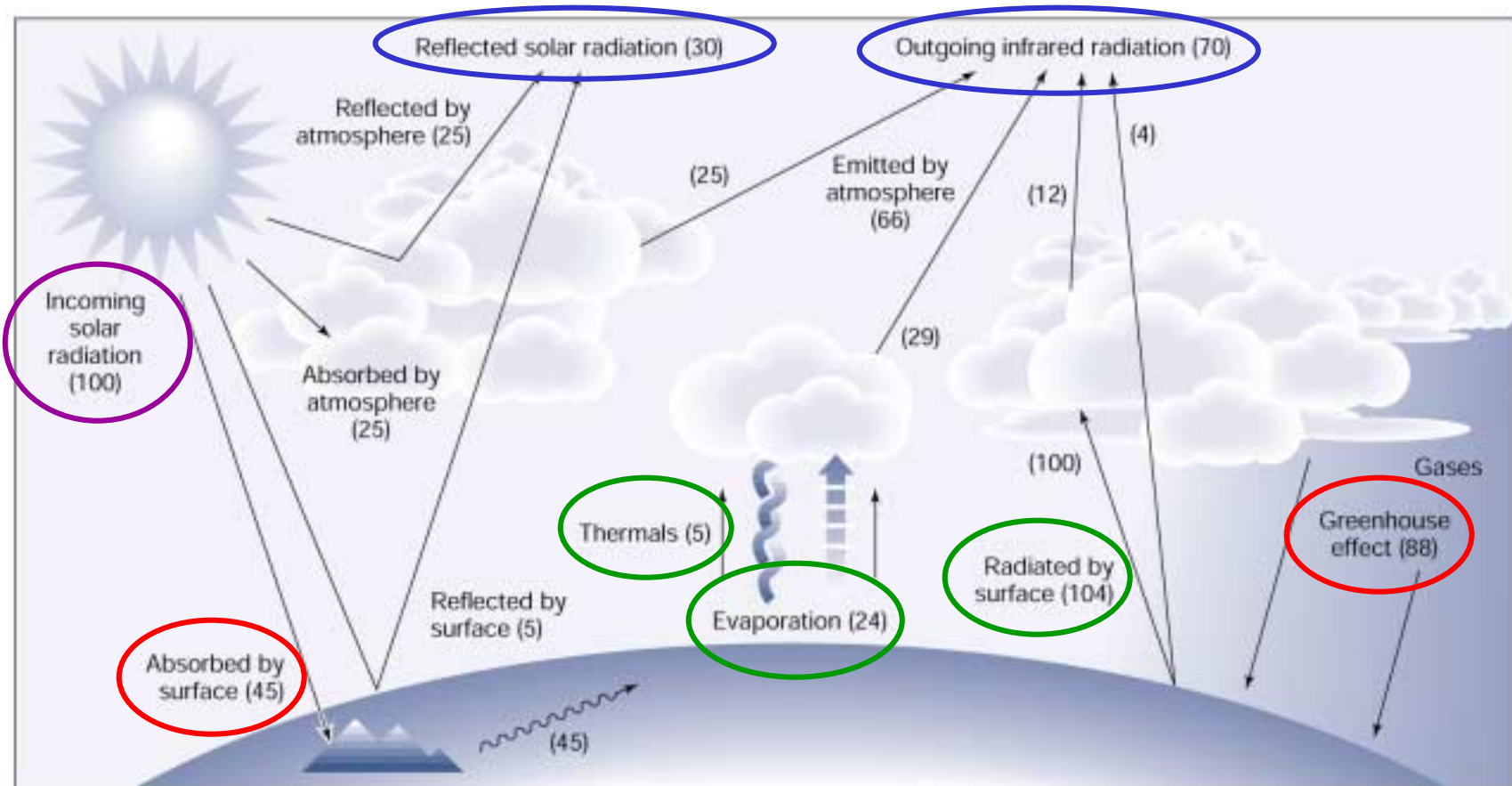
Mars  
(virtually no  
greenhouse)



Note: formula from the Text Box on page 43 gives wildly inaccurate estimates of greenhouse effect for Venus and Mars.

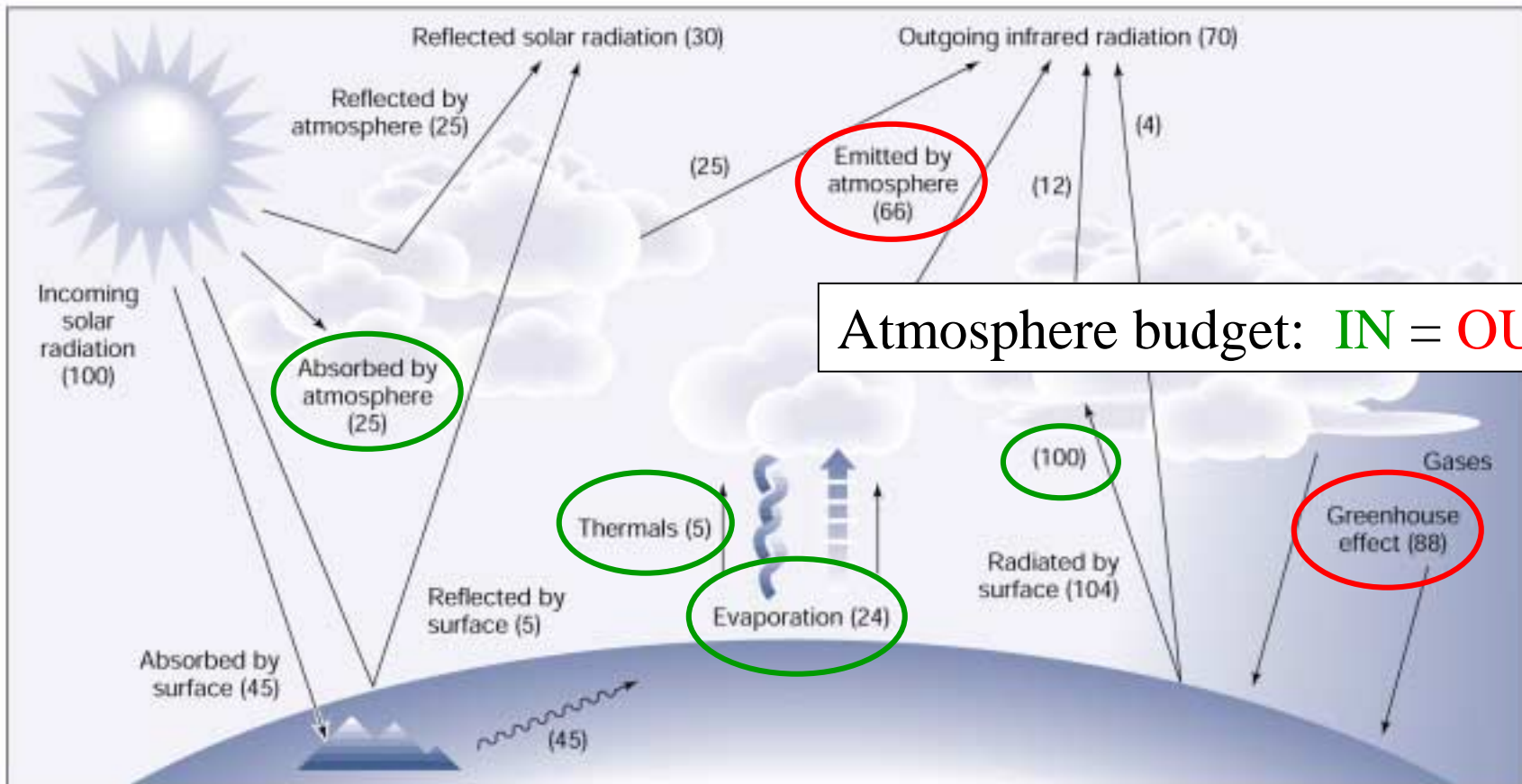
# Energy Budget: TOA and surface

Top-of-Atmosphere budget: **IN = OUT**



Surface budget: **IN = OUT**

# Energy Budget: atmosphere



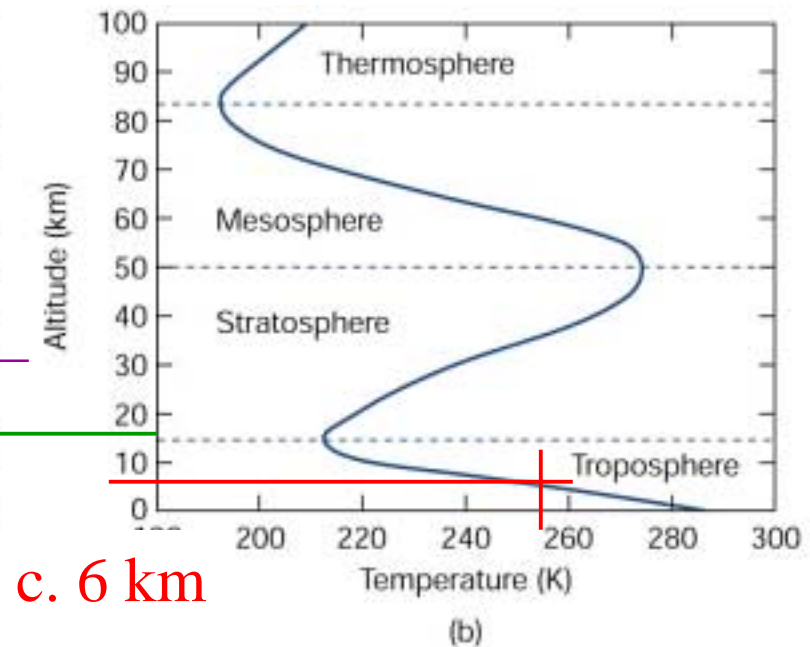
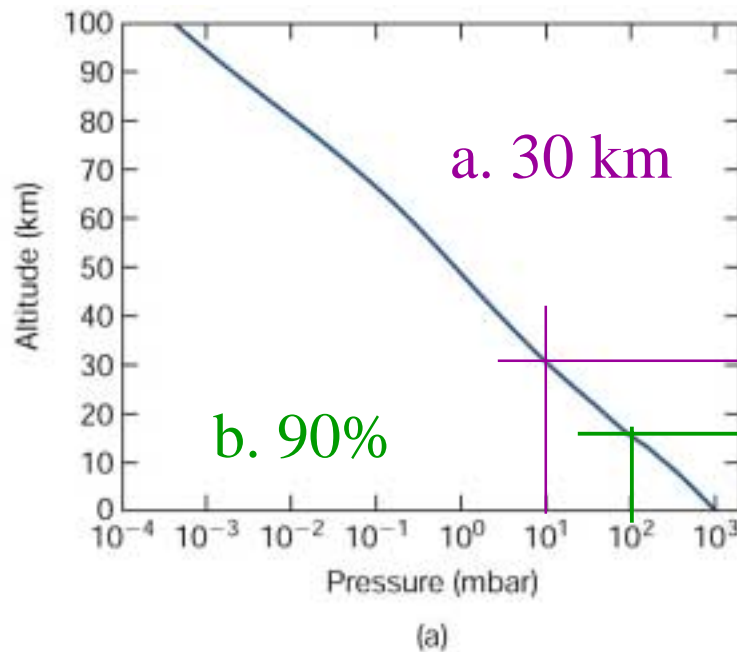
## Vertical structure of the atmosphere

Pressure: the weight of air overhead (e.g. psi or lbs/in<sup>2</sup>)

Note log scale on pressure graph.

a. At what height are you above 99% of the atmospheric mass?

b. What fraction of atmospheric mass is in the Troposphere?



Lapse rate: The rate at which temperature changes with height in the atmosphere. A measure of vertical stability.

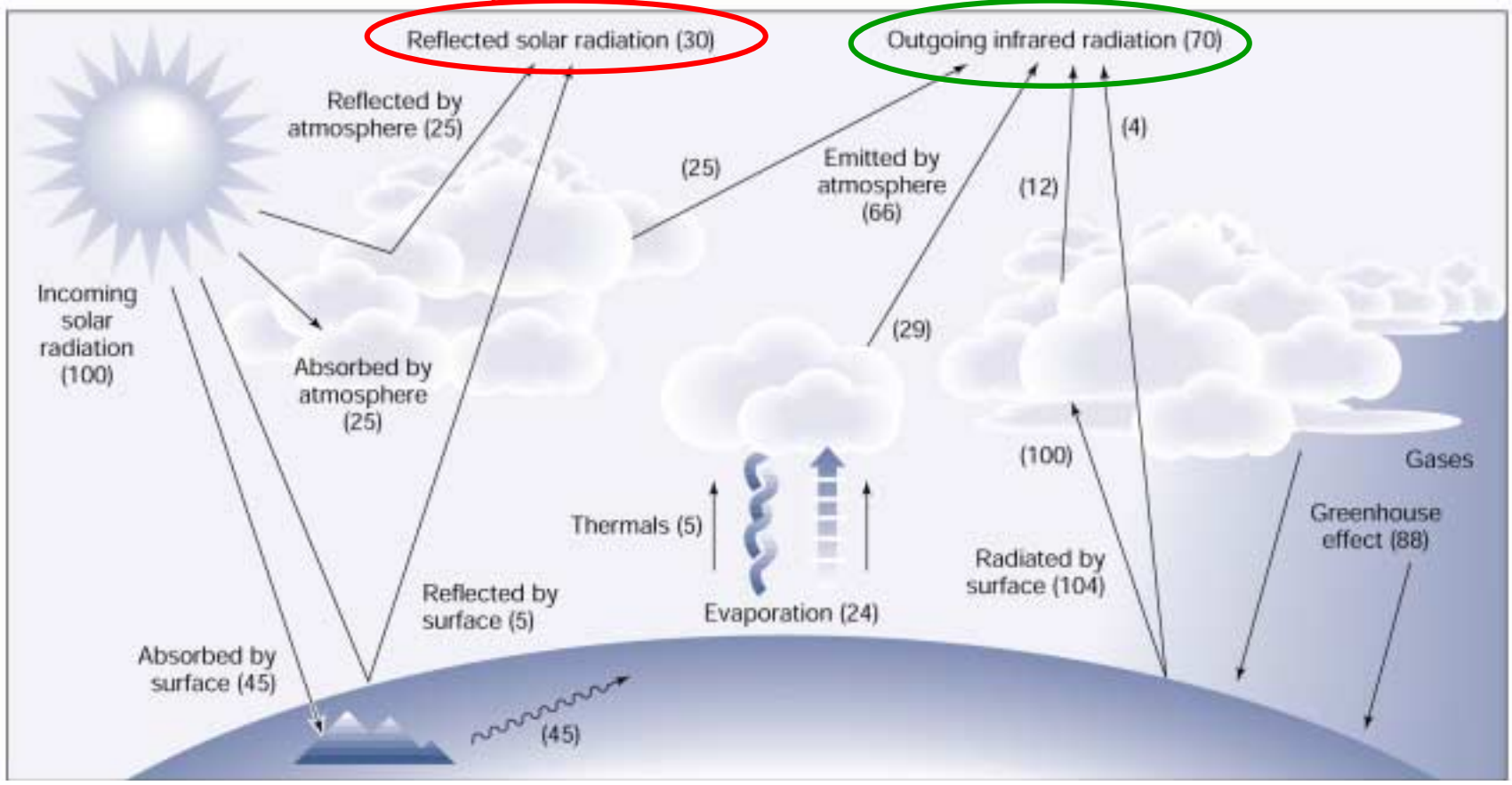
c. What height corresponds to  $T_e$ ?



# Satellite View of Energy Budget

visible satellite

infrared satellite





Wed Oct 22

Announcements:

- HW#2 due Thursday at beginning of class
- Monday's record rain

Today:

- vertical structure, "lapse rate"
- atmospheric composition
- greenhouse gas absorption - how does it work?
- dual role of clouds

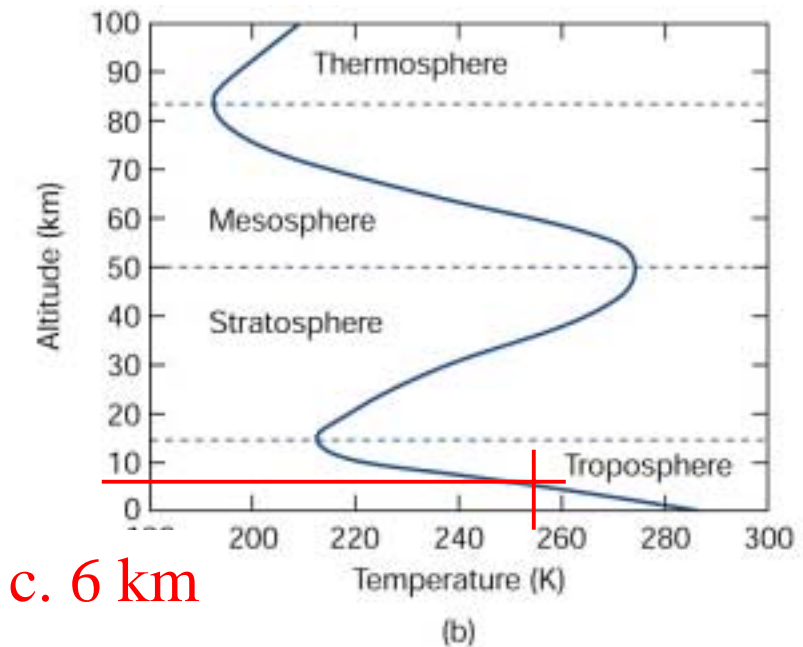
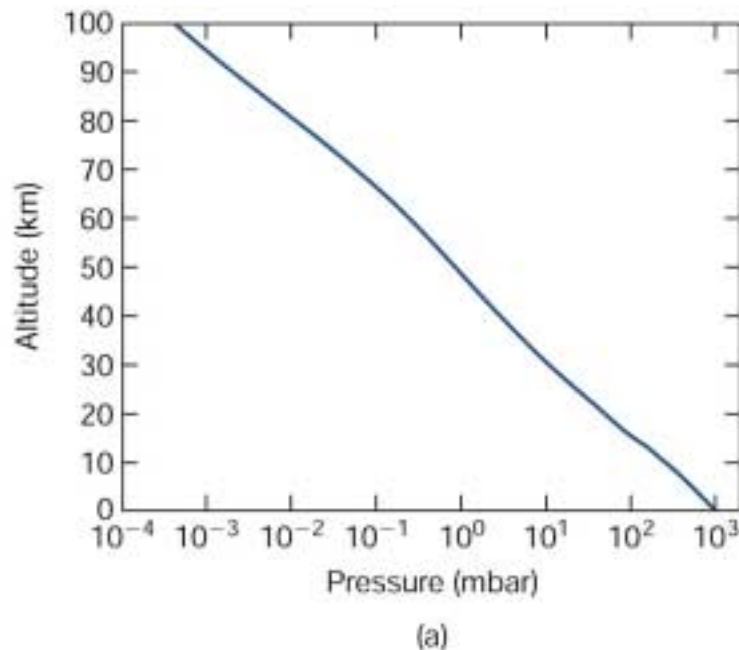
Speaking of atoms...

## Vertical structure of the atmosphere

Pressure: the weight of air overhead (e.g. psi or lbs/in<sup>2</sup>)

a. At sea level,  $P = 15$  psi. Why don't we feel it?

b. At what depth underwater is pressure doubled?



c. 6 km

Lapse rate: The rate at which temperature changes with height in the atmosphere. A measure of vertical stability.

c. What height corresponds to  $T_e$ ?

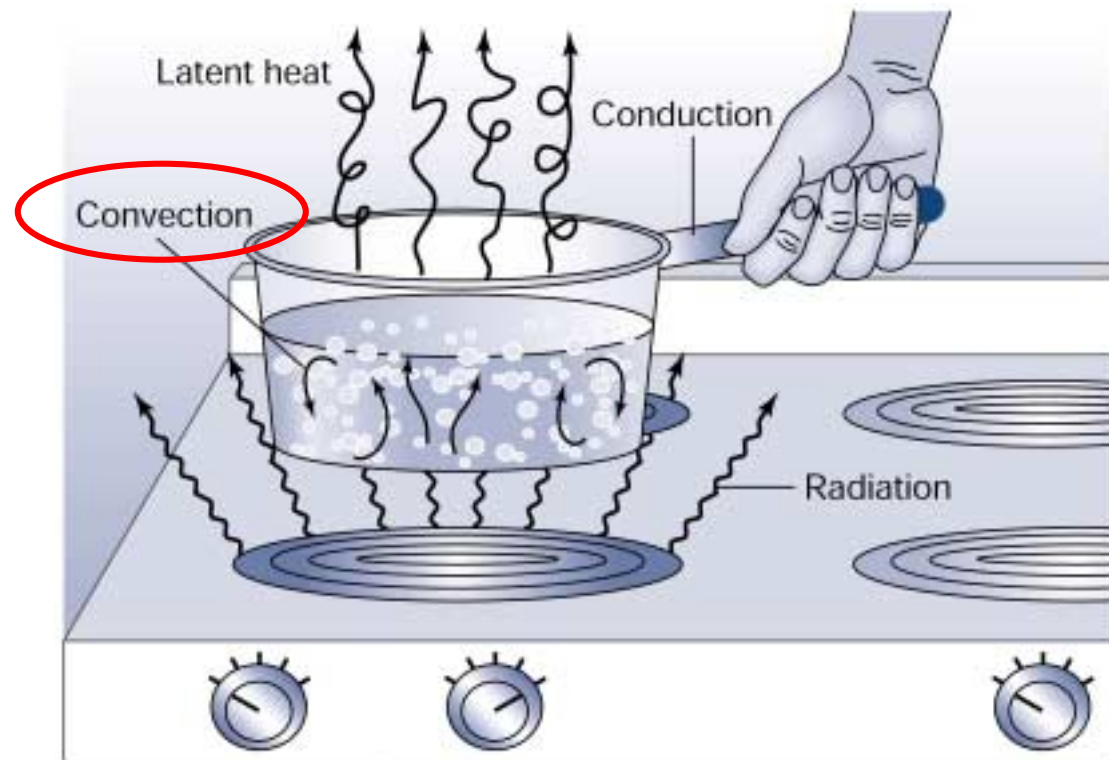
## Lapse rate and convection

Lapse rate: The rate at which temperature changes with height in the atmosphere. A measure of vertical stability.

Stable: Warm fluid on top of cold fluid.

Unstable: Warm fluid below cold fluid.

Box Fig. 3-3



## A dialectic on atmospheric vertical structure

Weather (clouds, rain) happens in the Troposphere because the Troposphere is unstable.

Why is the troposphere unstable?

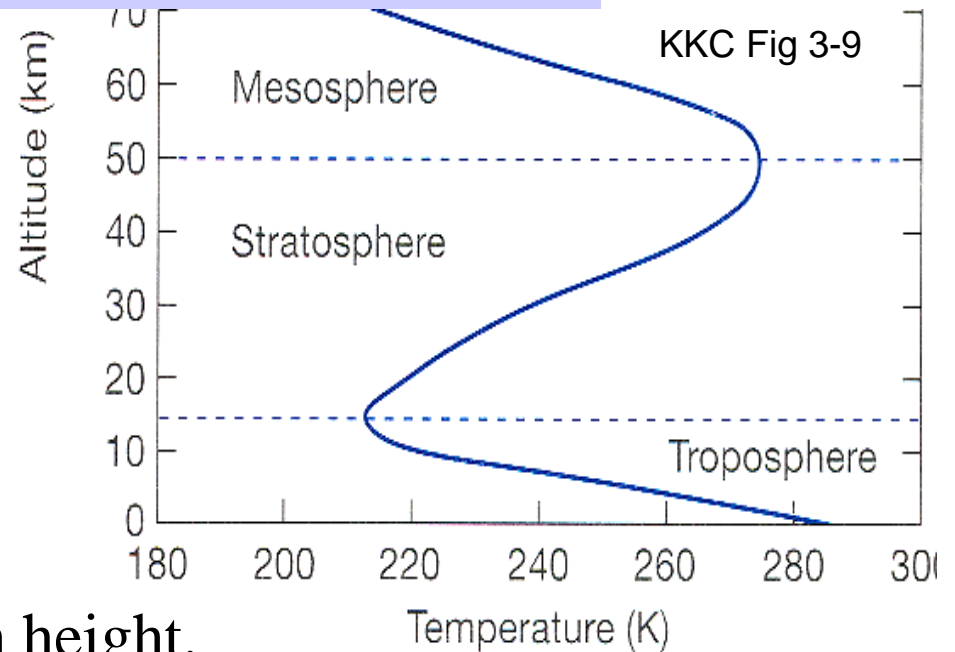
Because temperature decreases with height.

Yes, but why does temperature decrease with height?

Because it is heated from below.

Yes, but how does it come about that it is heated from below?

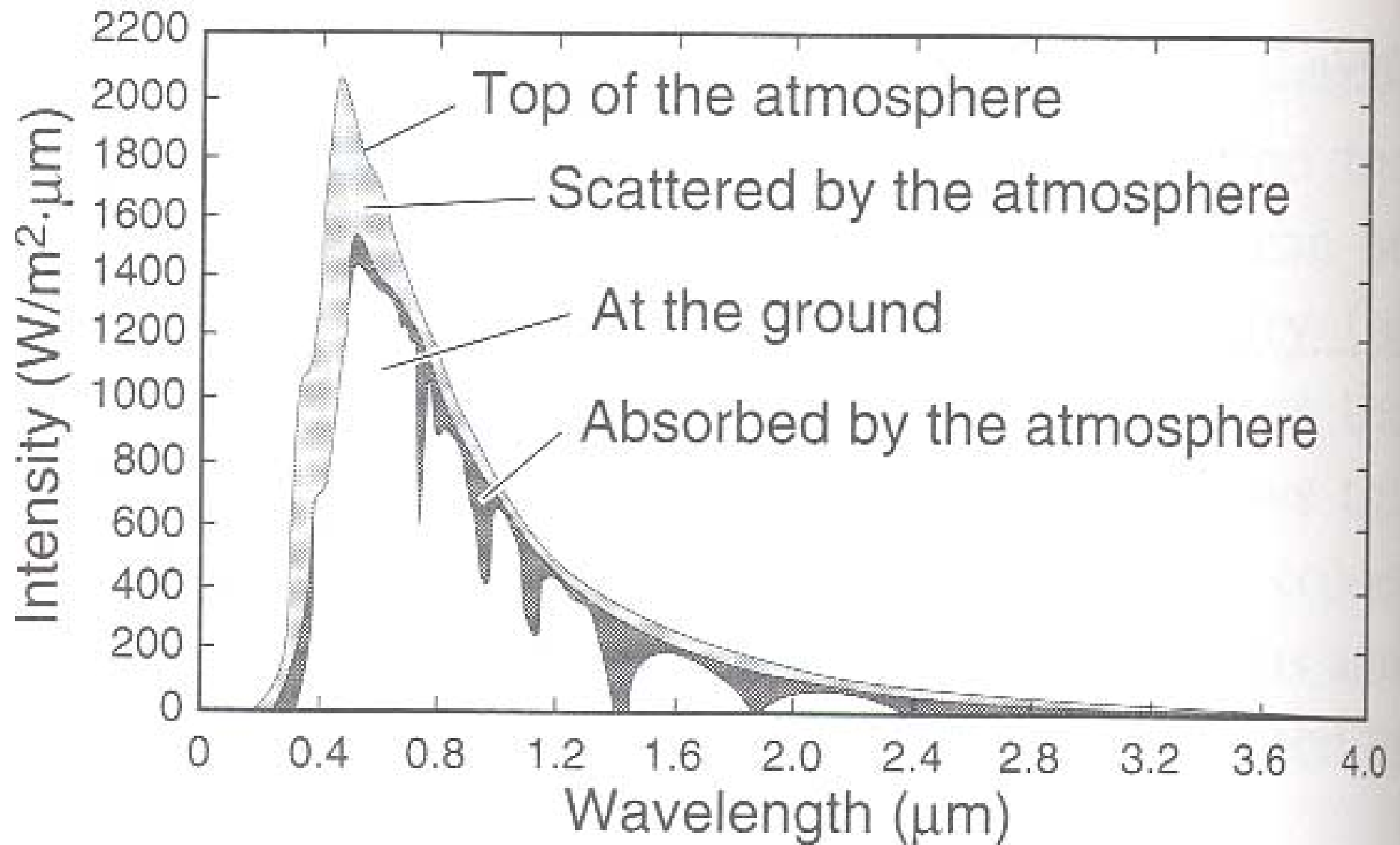
Because the atmosphere is mostly transparent to solar (or, SW) radiation.



## spectral absorption by the atmosphere

**SW**

most (non-reflected) SW radiation passes through the atmosphere

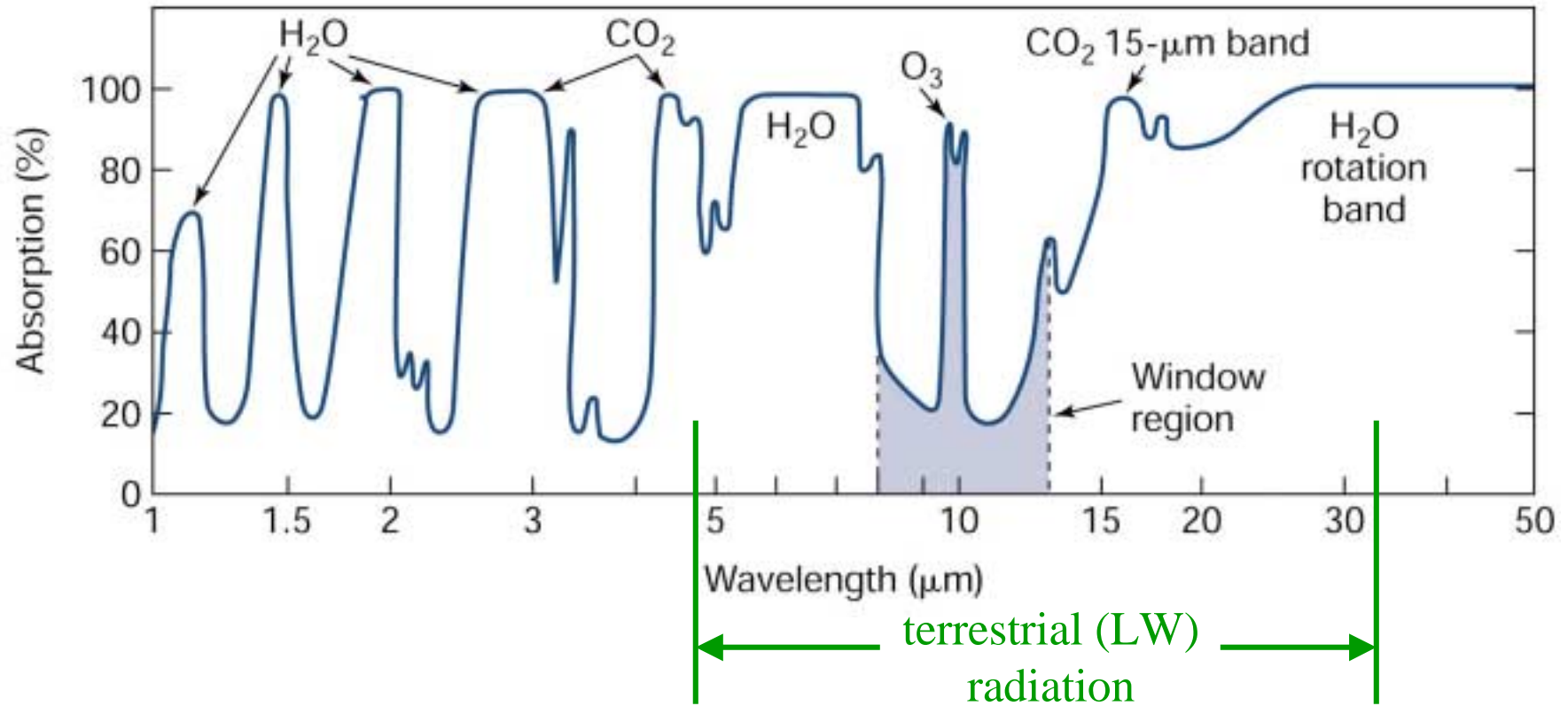


# spectral absorption by the atmosphere

most LW radiation emitted by the surface is absorbed by the atmosphere

LW

Fig 3-13



## Atmospheric constituents

- **main gases** ( $O_2$ ,  $N_2$ , Ar are 99.9% of dry volume of atmosphere)
- **variable gas** (water vapor)
- **trace gases** ( $CO_2$ ,  $O_3$ , CFC's,  $SO_2$ , etc, etc, etc)
- **particles** (dust, seasalt, sulfate, soot, organics)

**TABLE 3-2**

<b>Major Constituents of Earth's Atmosphere Today</b>	
<i>Name and Chemical Symbol</i>	<i>Concentration (% by volume)</i>
Nitrogen, $N_2$	78
Oxygen, $O_2$	21
Argon, Ar	0.9
Water vapor, $H_2O$	0.00001 (South Pole)–4 (tropics)
Carbon dioxide, $CO_2$	0.037*

\*In 2002

## Greenhouse gases

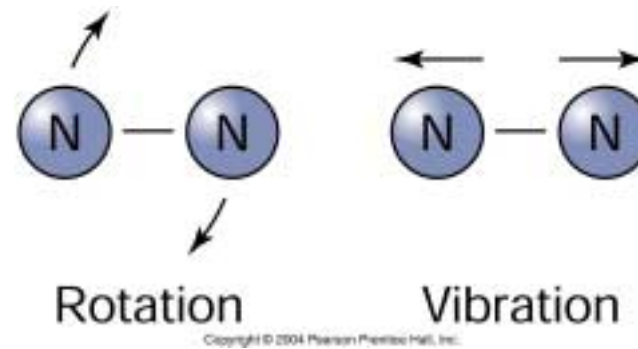
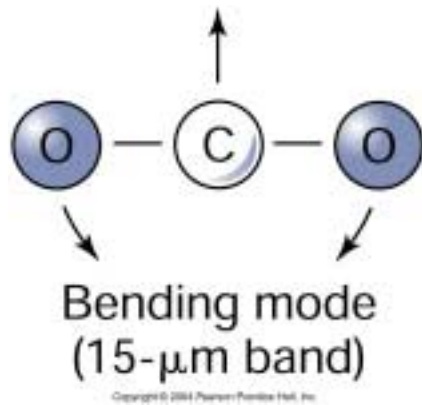
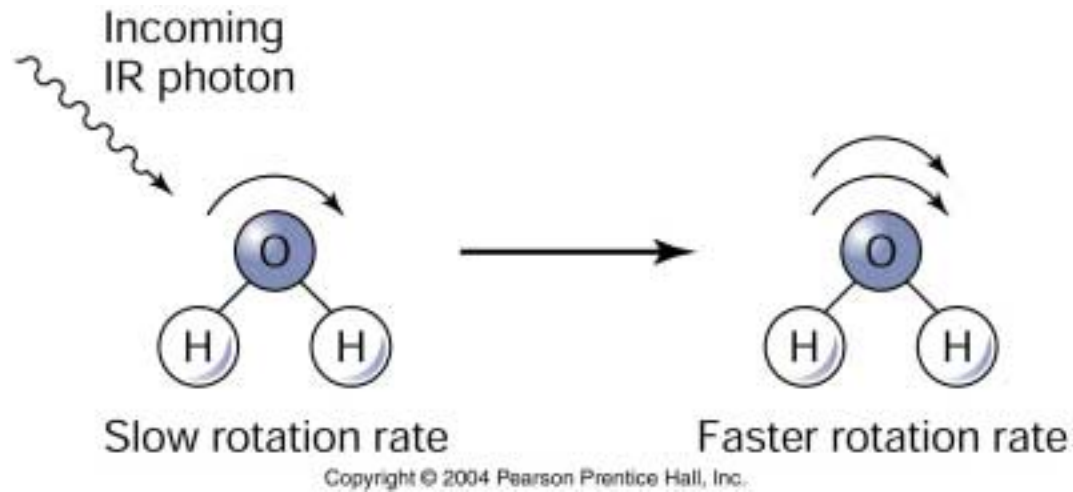
ppmv, parts per million by volume:  
molecules of trace substance per million molecules of air

**TABLE 3-3**

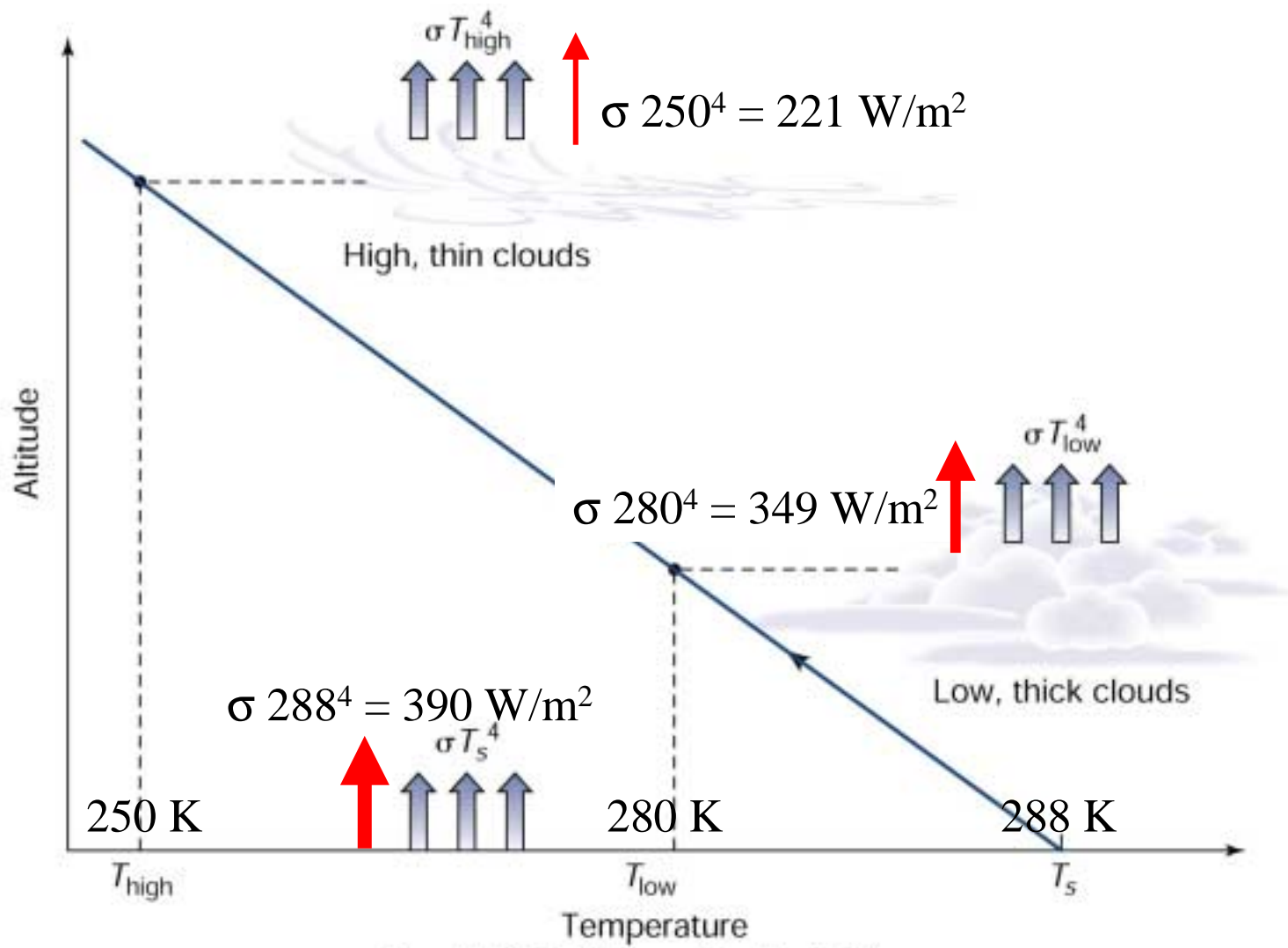
<b>Important Atmospheric Greenhouse Gases</b>	
<i>Name and Chemical Symbol</i>	<i>Concentration (ppm by volume)</i>
Water vapor, H <sub>2</sub> O	0.1 (South Pole)–40,000 (tropics)
Carbon dioxide, CO <sub>2</sub>	370
Methane, CH <sub>4</sub>	1.7
Nitrous oxide, N <sub>2</sub> O	0.3
Ozone, O <sub>3</sub>	0.01 (at the surface)
Freon-11, CCl <sub>3</sub> F	0.00026
Freon-12, CCl <sub>2</sub> F <sub>2</sub>	0.00054



# GHG absorption



# hi-lo cloud: Bad Figure!



## upcoming talks

FRIDAY 24 October:

**3:30 14 OTB (Oceanography Teaching Bldg)**

Dr. Ralph Keeling

Scripps Institute of Oceanography

"Fate of anthropogenic CO<sub>2</sub> and changing biogeochemistry  
of the oceans"

Thur Oct 23

### Announcements:

- HW #2 due today
- Readings for next week are on the web
  - Modeling chapter from new edition of book
  - Lorius et al. article on ice-ages and climate sensitivity
- midterm next Friday
- paper #1 due next Friday

### Today:

- multiple role of clouds in energy balance
- local energy balance, diurnal effects

### Tomorrow (Friday):

- guidance on papers
- review/questions
- supplementary topics

# Heat-trapping by clouds

calculate the heat-trapping (W/m<sup>2</sup>) by the cloud

Cloud:

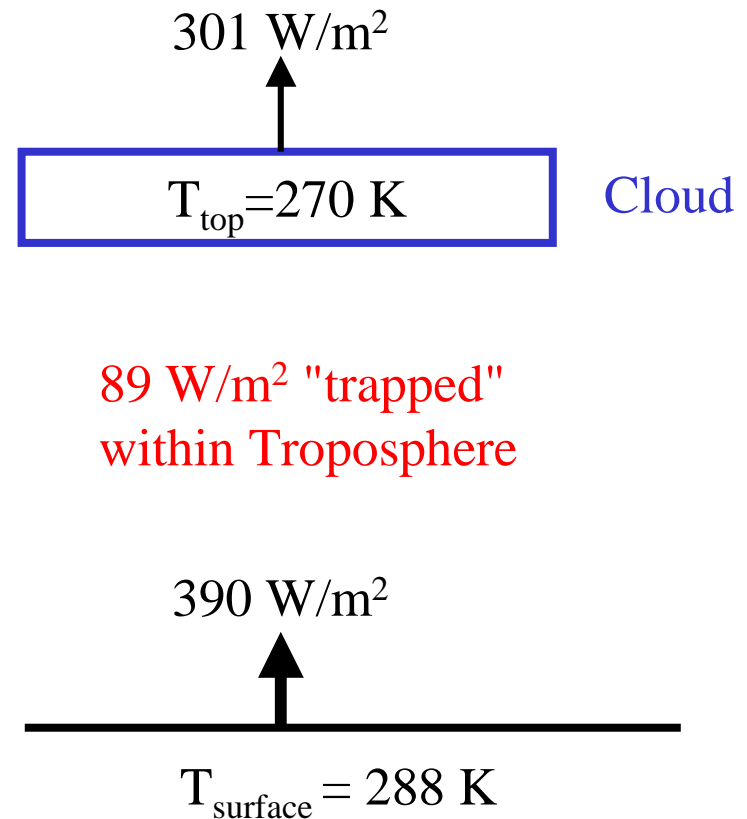
$$\begin{aligned} E_{\text{OUT}} &= \sigma T^4 \\ &= 5.67\text{e-}8 * 270^4 \\ &= 301 \text{ W/m}^2 \end{aligned}$$

Surface:

$$\begin{aligned} E_{\text{OUT}} &= \sigma T^4 \\ &= 5.67\text{e-}8 * 288^4 \\ &= 390 \text{ W/m}^2 \end{aligned}$$

Heat-trapping

$$\begin{aligned} &= 390 - 301 \\ &= 89 \text{ W/m}^2 \end{aligned}$$



What assumptions underlie this calculation?

- All LW radiation is absorbed by cloud.
- Cloud emits according to cloud-top temperature.

## Greenhouse role of clouds

**LW**

The previous slide demonstrates that clouds are a major participant in the Earth's greenhouse effect.

warming

**SW**

But...

We know from previous lectures that clouds are a major player in controlling Earth's albedo.

cooling

**net**

Question...

What is the net effect of a cloud on the energy balance?

???

## Local Energy Balance

calculate  $E_{\text{IN}}$ ,  $E_{\text{OUT}}$ , and net energy effect for each cloud type

High thin cloud

$$T=250 \text{ K} \quad A=0.20$$

$$\begin{aligned} E_{\text{IN}} &= S_0/4(1-A) \\ &= 1370/4*0.8 \\ &= 274 \text{ W/m}^2 \end{aligned}$$

$$\begin{aligned} E_{\text{OUT}} &= \sigma T^4 \\ &= 5.67\text{e-}8*250^4 \\ &= 221 \text{ W/m}^2 \end{aligned}$$

$$\begin{aligned} E_{\text{NET}} &= E_{\text{IN}} - E_{\text{OUT}} \\ &= 274 - 221 \\ &= +53 \text{ W/m}^2 \end{aligned}$$

$$\begin{aligned} E_{\text{IN}} &= S_0/4(1-A) \\ &= 1370/4*0.4 \\ &= 137 \text{ W/m}^2 \end{aligned}$$

$$\begin{aligned} E_{\text{OUT}} &= \sigma T^4 \\ &= 5.67\text{e-}8*280^4 \\ &= 349 \text{ W/m}^2 \end{aligned}$$

$$\begin{aligned} E_{\text{NET}} &= E_{\text{IN}} - E_{\text{OUT}} \\ &= 137 - 349 \\ &= -212 \text{ W/m}^2 \end{aligned}$$

$$T=280 \text{ K} \quad A=0.60$$

Low, thick cloud

# Multiple roles of clouds in the energy budget

SW: Clouds are the major player in the Earth's albedo

LW: Clouds are a major player in heat-trapping (greenhouse effect)

high clouds:

modest albedo (small SW effect)

lots of "heat-trapping"

complete IR absorption

cold, so have low IR emission

low clouds:

high albedo (big SW effect)

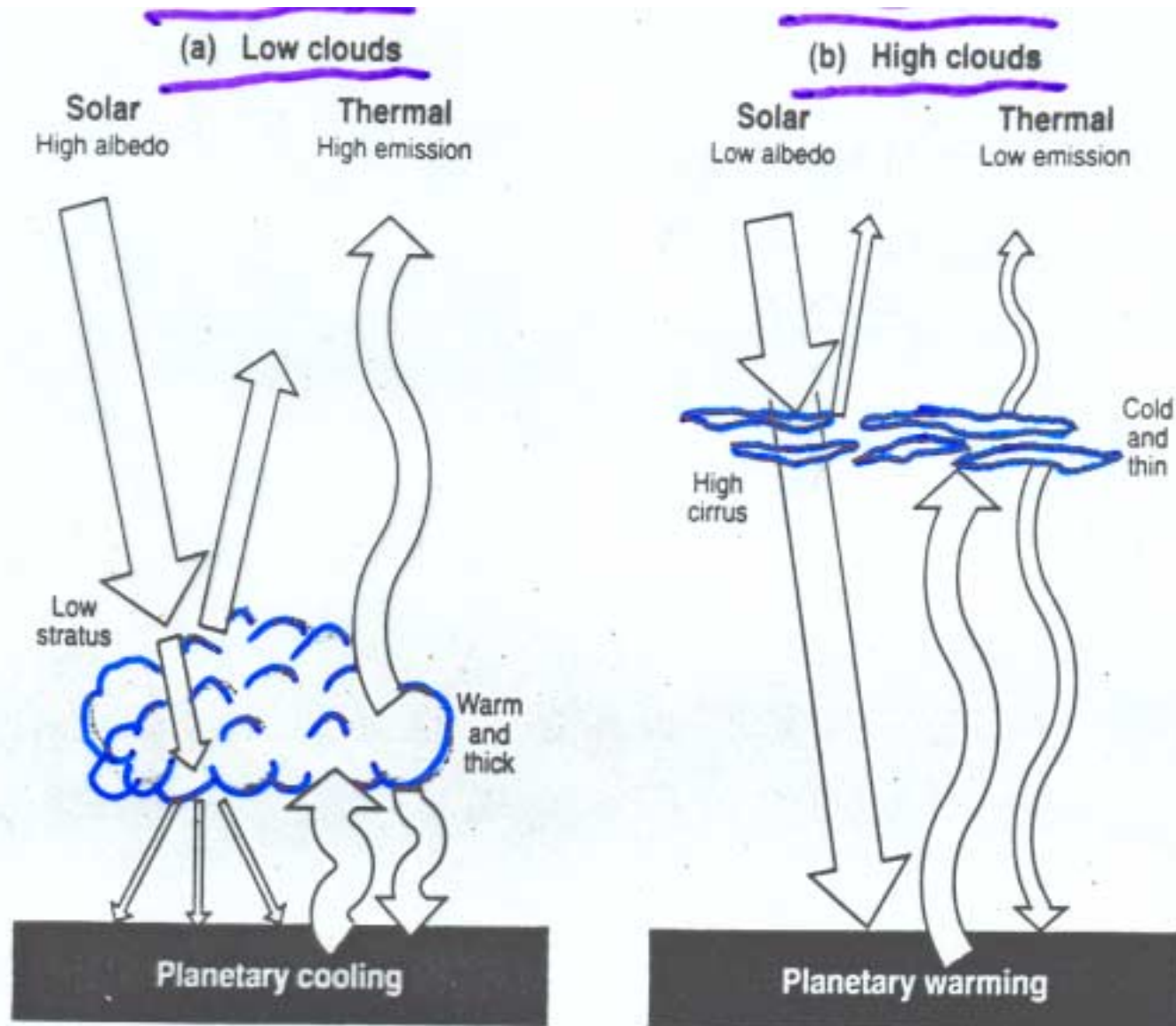
modest "heat-trapping"

complete IR absorption

almost as warm as the surface, so emit almost  
as much IR upwards



# High vs Low Clouds graphic



# Home Weather record

Where on these plots do you see the SW effect of clouds?

Where do you see the LW effect of clouds?

